

Palynostratigraphical and oceanic anoxic events investigation in the Lower Jurassic Black Shale, Indus Basin, Pakistan

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Abstract

The stratigraphy of Lower Jurassic in the Indus Basin, Pakistan, portrays the siliciclastic and carbonate strata in studied sections. The representative stratigraphic sections were sampled to investigate using carbon and oxygen isotopes ($\delta^{13}\text{C}$ & $\delta^{18}\text{O}$), and palynology. Based on diagnostic palynomorphs from the Lower Jurassic, a single Hettigian-Pliensbachian assemblage biozone (HPABZ) is recognized. The oceanic anoxic event with a 2–3 ‰ negative shift in $\delta^{13}\text{C}$ & $\delta^{18}\text{O}$ across the Pliensbachian-Toarcian boundary is noted in the studied sections. The nature of the oceanic anoxic events are discussed and correlated with the established events worldwide.

Keywords: Black shale; Jurassic; Palynology; Isotopes; Indus Basin, Pakistan.

1. Introduction

Palynostratigraphy is widely used to establish reliable biostratigraphy and to determine depositional environments (Weerakoon et al., 2019). Palynomorphs serve as authentic proxies for precise and reliable paleo-reconstructions of paleoecology, paleogeography and paleoclimatology (Césari and Colombi, 2016; Lindström et al., 2016). However, in case of the Indus Basin very limited research has been done on the palynostratigraphy and stable isotope stratigraphy of the Jurassic strata in the studied area. The previous studies have addressed the sedimentological and stratigraphic aspects of Lower to Middle Jurassic succession based on lithofacies, microfacies and petrography (Fatmi, 1974; Akhtar, 1983; Mensink et al., 1988; Fatmi et al., 1990; Mertmann and Ahmad, 1994; Ali et al., 2018; 2019; Iqbal et al., 2019; Saboor et al., 2022; Ullah et al., 2022a). Only few reports of palynostratigraphy are available for this succession (Masood et al., 1997; Ali et al., 2018). The Jurassic (201–145 Myr ago) was a warm greenhouse period. Its warmest interval is represented by the late Early Jurassic Toarcian Stage, which is characterized by an unexpected ~7°C increase globally at the beginning of Toarcian oceanic anoxic events (TOAE) in the mid-latitudes (Dera et al., 2011;

Gradstein et al., 2012; Korte et al., 2015). Previous work on the high-resolution carbon-isotope data ($\delta^{13}\text{C}$) on Lower to Middle Jurassic periods, especially the Pliensbachian-Toarcian boundary (PL-TO) is scarce to non-existent in the Indus Basin exist as compared to the extensive studies elsewhere in the world (Hesselbo et al., 2000; Jenkyns et al., 2001, 2002; Katz et al., 2005; McElwain et al., 2005; Svensen et al., 2007; Littler et al., 2010; Kemp et al., 2011; Kemp and Izumi, 2014; Ali et al., 2018; 2022 a,b Izumi et al., 2018; Xu et al., 2018). The conventional stratigraphy and sedimentology of Indus Basin have been addressed in many of previous studies but no formal biostratigraphic framework and study of the anoxic events has been established yet in the Jurassic clastic and carbonate successions. In order to establish a reliable stratigraphy of any rock unit, it is significant to first put the same rock unit into a reliable biostratigraphic framework. The age of the Datta Formation is based on the law of superposition as no age related index fossil has been reported from it (Shah, 2009). The Shinawari Formation overlies the Datta Formation and contains some Toarcian ammonites, so the age assigned to the Datta Formation is based on stratigraphic assumptions (Shah, 2009). Therefore, this study is focused on establishing a reliable stratigraphic framework, augmented by stable

isotope stratigraphic analysis, to demarcate age as well as the oceanic anoxic events in studied strata.

2. Geological Setting

After rifting from the Eastern Gondwana landmass, the Indian Plate started drifting north during the Early Cretaceous (c. 130 Ma), and collided with the Eurasian Plate giving rise to the Himalayas (Coward et al., 1986; Gibbons et al., 2013). The drifting of the Indian Plate squashed the Neo-Tethys Ocean, present between India and Eurasia, causing intra-oceanic subduction which subsequently gave rise to the Kohistan Island Arc (Searle et al., 1987; Searle, 1991). The Kohistan Island Arc (KIA) is dominantly composed of plutonic, volcanic and metasedimentary rocks covering an area of about 36000 km². The KIA collided with the Karakoram Block during the Cretaceous along the suture zone known as Main Karakoram Thrust (MKT) (Windley, 1983). Later, in the Cenozoic (c. 55-50 Ma) the Indian Plate subducted below the Kohistan Island Arc and formed the Main Mantle Thrust

(MMT), also known as the Indus Suture Zone (Tahirkheli, 1982). After the formation of MMT, the deformational stresses shifted southward and formed many regional thrusts, such as Main Boundary Thrust (MBT) and Salt Range Thrust (SRT). The Indus Basin is bounded by MBT in the north and Arabian Sea in the south and Chaman Transform Fault in the northwest. The Indus Basin preserves excellent sedimentary successions from Pre-Cambrian to Recent. The Pre-Cambrian and Paleozoic strata is only exposed in the Lower Indus Basin whereas the Mesozoic and Cenozoic rocks are exposed throughout the Indus Basin (Shah, 2009; Babar et al., 2018; Zahir et al., 2021; Ullah et al., 2022b; Sabba et al., 2023). The Indus Basin Jurassic sediments were deposited in the eastern part of the Tethys Ocean, representing the northwestern margin of the Indian Plate. The Jurassic sediments are represent by the clastic dominated Lower Jurassic Datta and Shirinab formations and the carbonate dominated Shinawari and Samana Suk formations and Chiltan Limestone (Ali et al., 2013, 2019) Fig. 1).

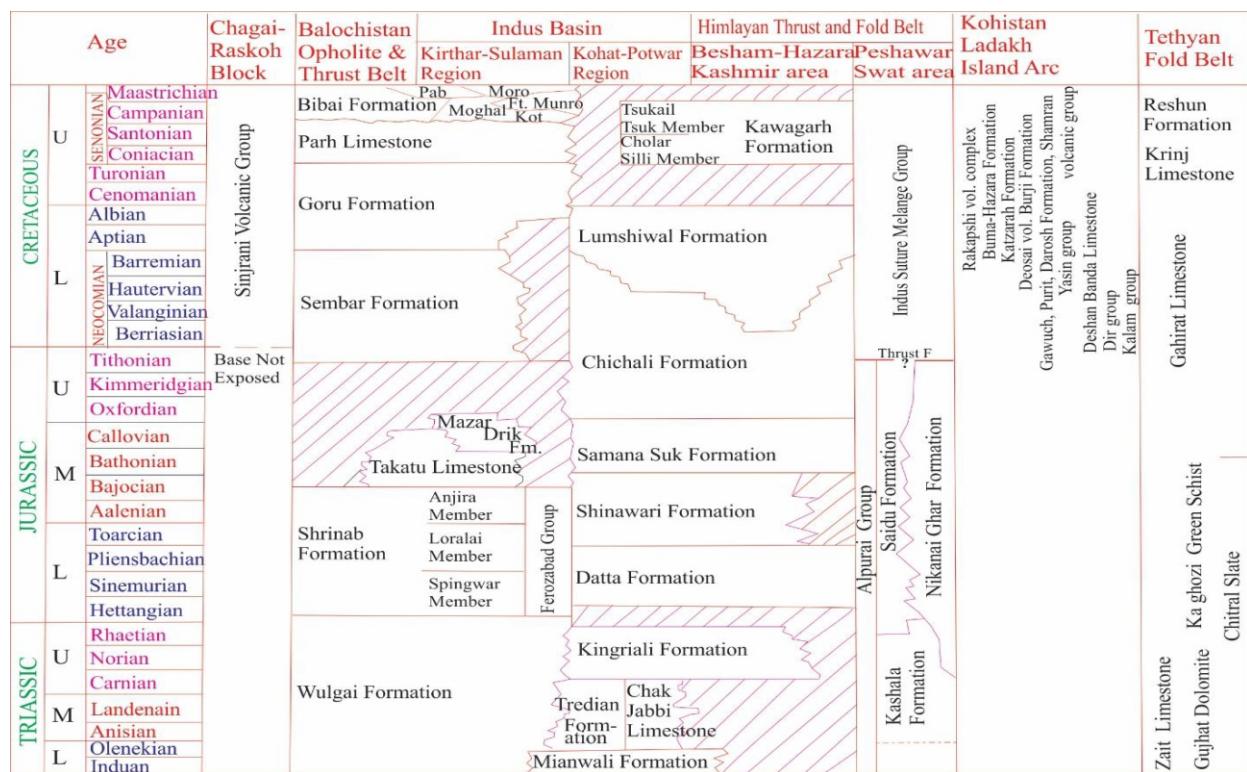


Fig. 1. Shows the Mesozoic stratigraphy of the Baluchistan Basin, Indus Basin and Kohistan-Ladakh Island Arc, Pakistan (modified after Bender and Raza, 1995).

3. Methods and Materials

A total of 29 fresh shale samples from outcrops of the Jurassic Datta and Shinawari formations from the following five sections 1. Kahi village; 2. Nammal Gorge Section; 3. Baroch Nala; 4. Chichali Nala; 5. Gula Khel Nala from Salt Ranges Pakistan were collected for this study (Fig. 2). The stratigraphic true thickness of the Datta and Shinawari formations were measured as per procedure described in previous published literature

(Prothero and Schwab, 2004), and field photographs were taken (Fig. 3). Palynological slides were prepared from the dark shales by following the standard procedure of maceration (Wood et al., 1996; Ali et al., 2018). Transmitted light microscopy (i.e. Nikon SMZ-25) and polarizing microscope (i.e. Nikon OS-F12) were used for palynological analysis. The Datta Formation provided a good quantity and quality of terrestrial palynomorphs, with some marine content, in all studied samples.

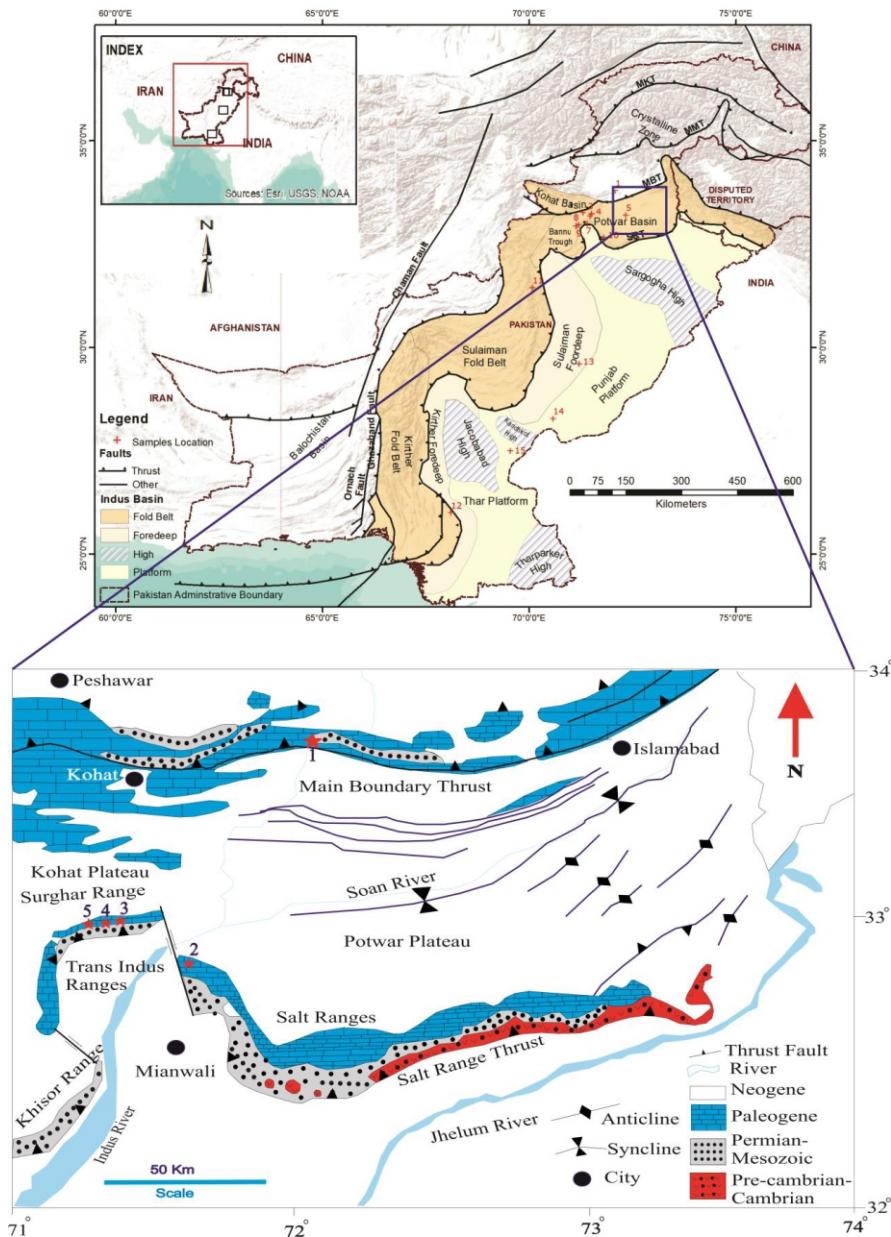


Fig. 2. Generalized geological map showing locations of the studied stratigraphic sections (denoted by steric and number) in the Indus Basin, Pakistan (used after Kazmi and Rana, 1982).
1. Kahi village; 2. Nammal Gorge Section; 3. Baroch Nala; 4. Chichali Nala; 5. Gula Khel Nala.

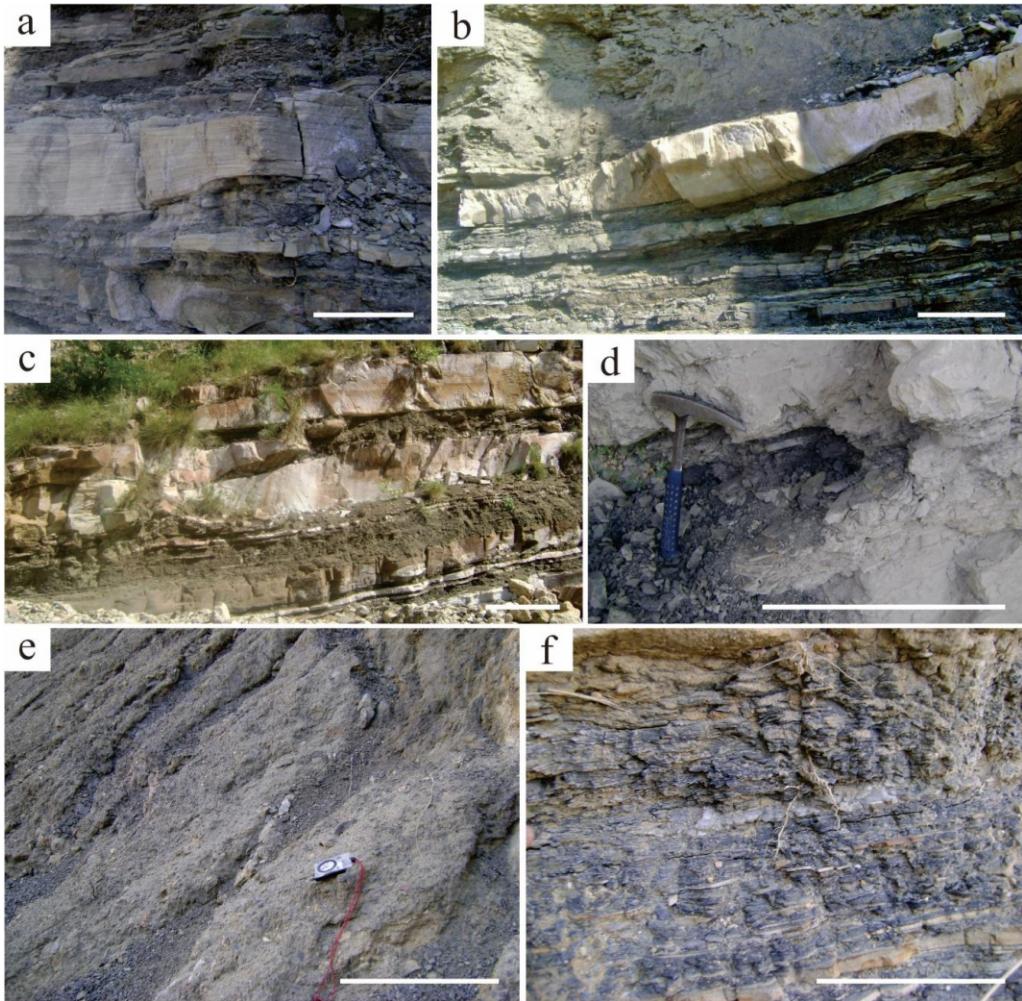


Fig. 3. Field photographs showing alternating beds of shale with minor limestone from the Jurassic Datta (a-c) and Shinawari (d-f) formations. (Scale: Length of the bar is 50cm)

The isotopic analysis for carbon and oxygen were carried out on a VG Isogas Prism III (Wolfson Laboratory, UK) and the Deltaplus Advantage (Thermo Fisher Scientific) was used in the current geological investigation. According to Ali et al. (2018), the PACS-2 standard was used for isotopic calibration, while Acetaniide was used for elemental composition. The standard deviation limit of each measurement for $\delta^{18}\text{O}$ is $\pm 0.2\text{\textperthousand}$ and for $\delta^{13}\text{C}$ is $\pm 0.1\text{\textperthousand}$ in the current study.

The total organic carbon is determined from the Total carbon analyzer (TOC). The TOC were measured as per the criteria described in Ali et al. (2022a). The TOC of all samples were measured in helium inert internal temperature. A 500 mg fine grained rock samples was subjected to 10% HCl (hydrochloric acid) to remove the inorganic carbon from the samples.

In order to remove the effect of HCl the samples were washed with distilled water and subsequently placed in oven to remove the moisture. Further the specimens were placed in clay crucible and analyzed in the TOC analyzer (ELTRA HELIOSL).

4. Results

4.1 Lower Jurassic Palynostratigraphy

The Lower Jurassic palynostratigraphic taxa is represented by the following assemblage that includes the following stratigraphically significant taxa: *Calamospora tener*, *Vitreisporites pallidus*, *Artrisporites minimus*, *Chasmatosporites elegans*, *Chasmatosporites hians*, *Chasmatosporites major*, *Corollina torosus*, *Corollina meyeriana*, *Cerebropollenites macroverrucosus*, *Alisporites pergrandis*, *Camptotriletes*,

cerebreformis, *Classopollis* sp., *Dipterella oblatinoides*, *Marattisporites scabrus*, *Osmundacidites* sp., *Stereisporites compactus*, *Stereisporites congregatus*, *Stereisporites*

bujargiensis, *Stereisporites psilatus* and *Stereisporites infragranulatus*, (Figs. 4-5 and Appendix . 1).

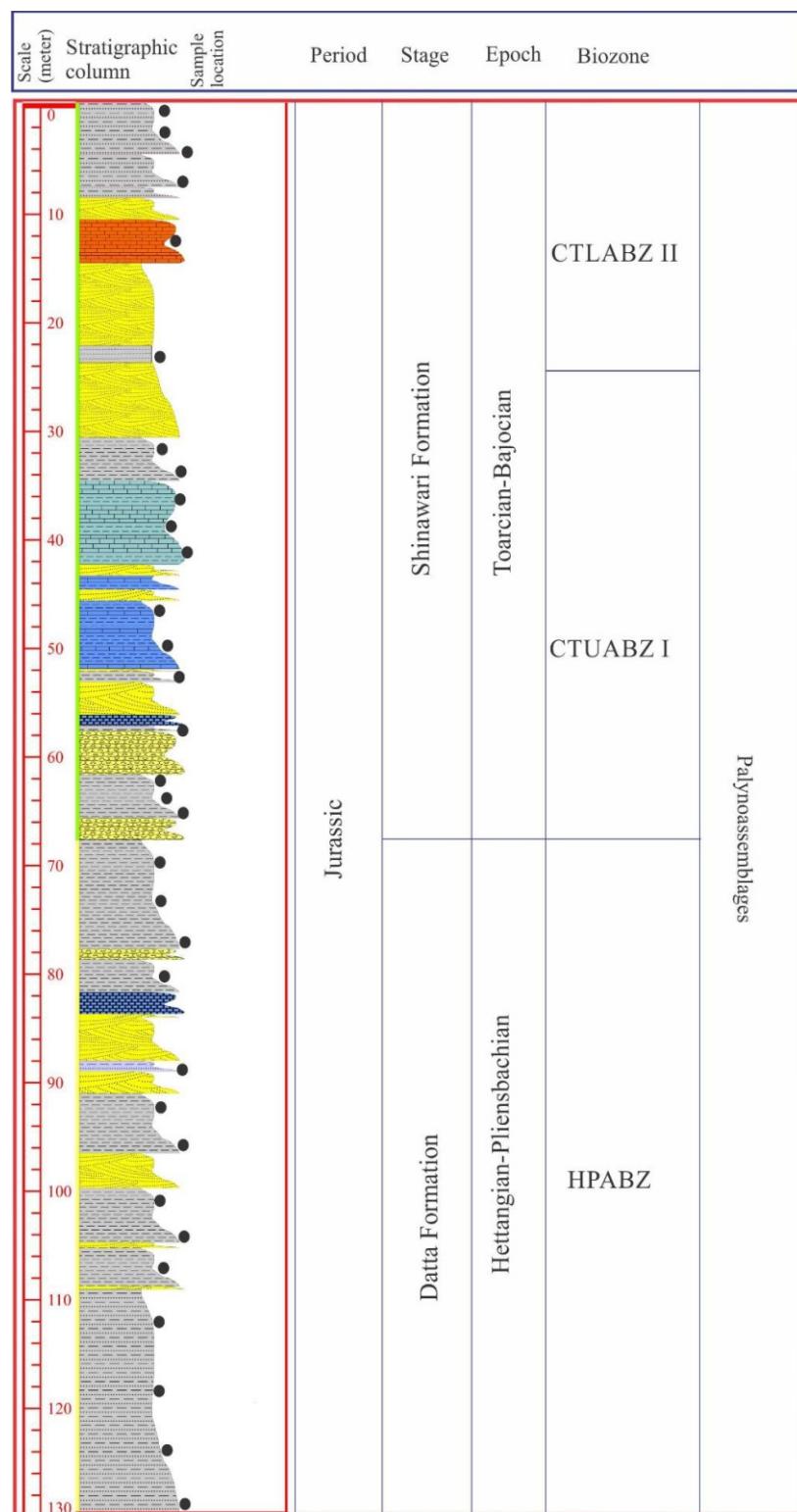


Fig. 4. Shows the stratigraphic column of the studied sections (Datta and Shinawari formations) and the assemblage biozones in the Indus Basin, Pakistan (the black circles represent sample locations).

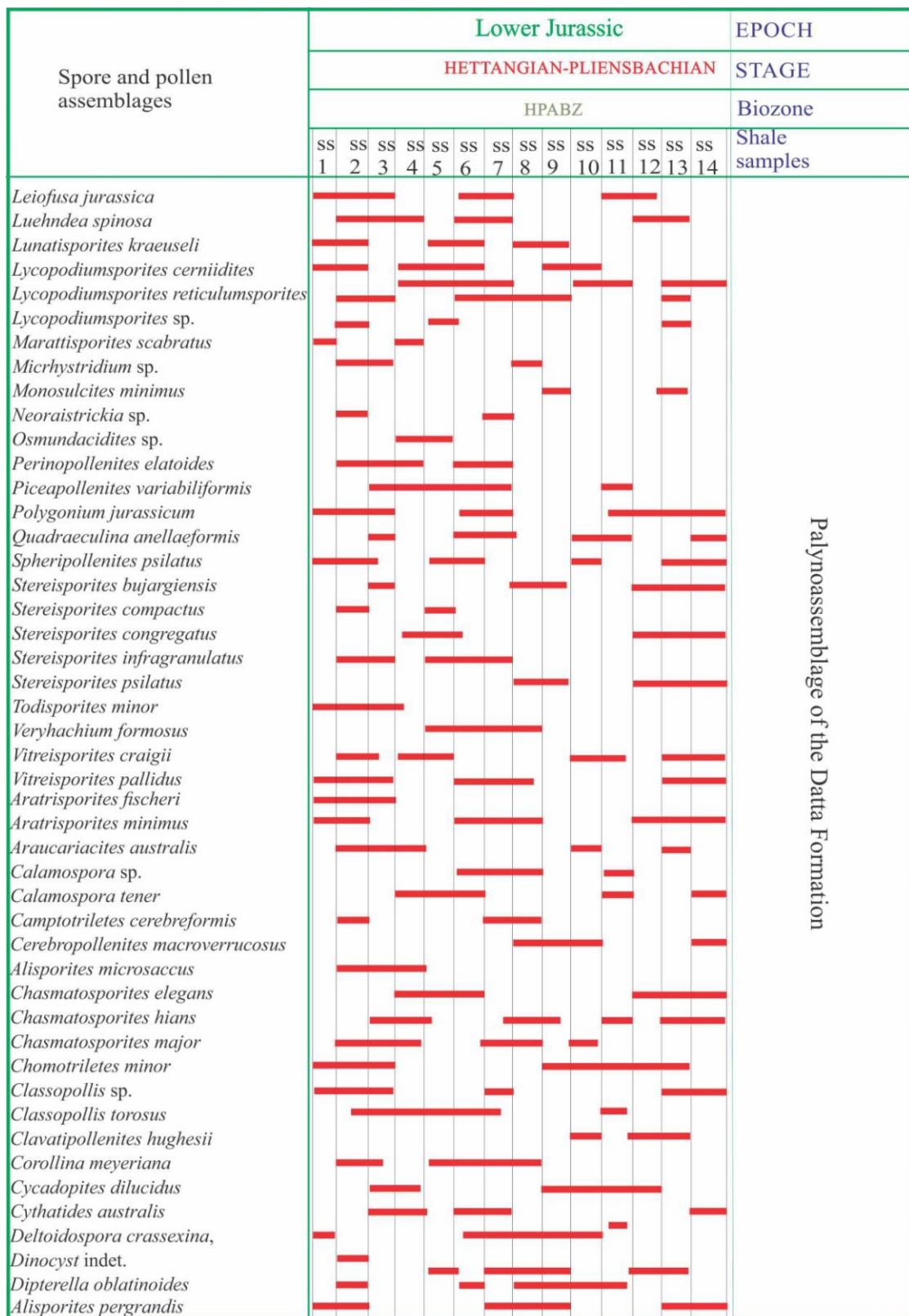


Fig. 5. Shows the range of different palyno-assemblage in the Lower Jurassic, Indus Basin, Pakistan.

Additional taxa include *Aratrisporites fischeri*, *Aratrisporites minimus*, *Araucariacites australis*, *Calamospora* sp., *Camptotriletes cerebreformis*, cf. *Alisporites microsaccus*, cf. *Chasmatosporites elegans*, *Chomotriletes minor*, *Corollina* sp., *Clavatipollenites hughesii*, *Cycadopites dilucidus*, *Cyathidites australis*, *Deltoidospora crassexina*, *Dinocyst* indet., *Leiofusa jurassica*, *Luehndea spinosa*, *Lunatisporites kraeuseli*, *Lycopodiumsporites cerniidites*, *Lycopodiumsporites*

reticulumsporites, *Lycopodiumsporites* sp., *Marattisporites* sp., *Micrhystridium* sp., *Monosulcites minimus*, *Neoraistrickia* sp., *Osmundacidites* sp., *Perinopollenites elatoides*, *Piceapollenites variabiliformis*, *Polygonum jurassicum*, *Pterospermella* sp., *Quadraceculina anellaeformis*, *Spheripollenites psilatus*, *Spheripollenites* sp., *Todisporites minor*, *Verrucosisporites* sp., *Veryhachium formosus*, and *Vitreisporites craigii* (Figs. 6-7).

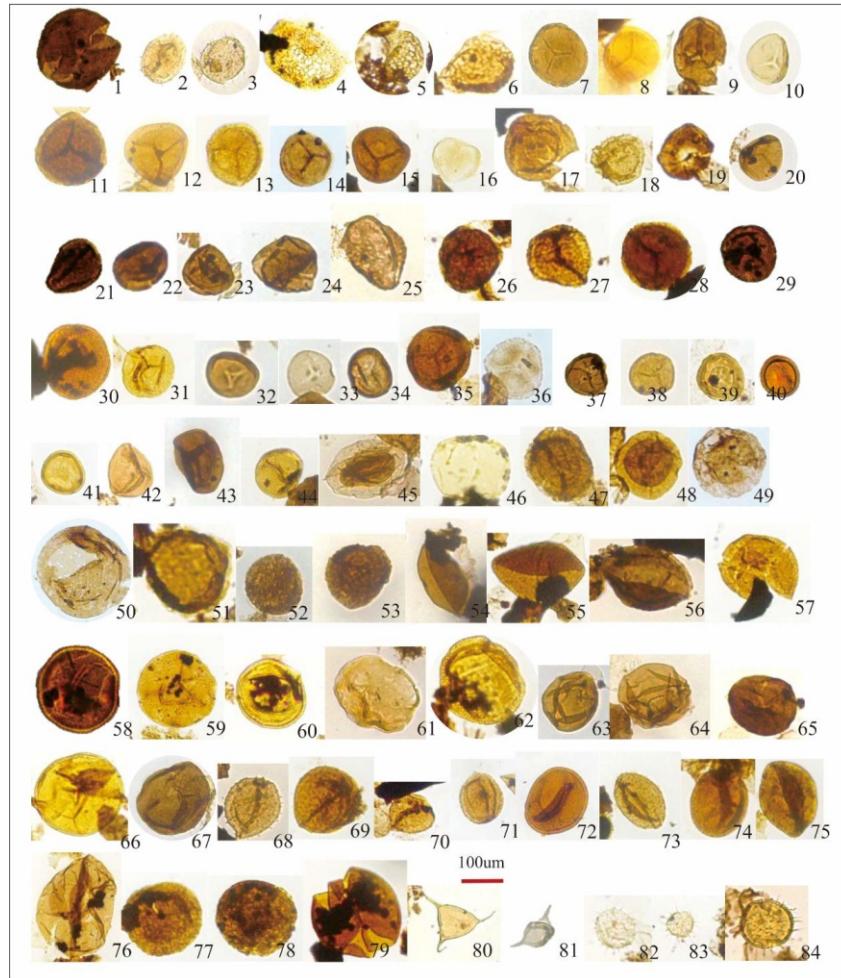


Fig. 6. Palynomorphs from the Lower Jurassic of the Indus Basin. 1. *Piceapollenites* sp.; Figs. 2-3. *Neoraistrickia* sp.; Figs 4-6. *Lycopodiumsporites* sp.; 7-8. *Stereisporites infragranulatus*; Figs 9-10. *Stereisporites congregates*; 11-15. *Stereisporites bujargiensis*; 16-19. *Classopollis torosus*; 20. *Pterospermella* sp.; 21-25. *Marattisporites scabratus*; 26-29. *Camptotriletes cerebreformis*; 30-31. *Osmundacidites* sp.; 32-33. *Stereisporites psilatus*; 34. *Stereisporites compactus*; 35-36. *Stereisporites infragranulatus*; 37-39. *Stereisporites congregates*; 40-421 *Exesipollenites laevigatus*; 42-44. *eltoidospora crassexina*; 45. *Chomotriletes minor*; 46. *Spheripollenites psilatus*; 47. *Clavatipollenites hughesii*; 48-49. *Corollina meyeriana*; 50. *Perinopollenites elatoides*; 51. *Lycopodiumsporites reticulumsporites*; 52-53. *Cerebropollenites macroverrucosus*; 54-55. *Marattisporites* sp.; 56. *onosulcites minimus*; 57. *Lycopodiumsporites cerniidites*; 58-61. *Todisporites minor*; 62. *Calamospora* sp.; 63-64. *Calamospora tener*; 65-76. *Aratrisporites minimus*; 77-78. *Aratrisporites* sp.; 79. *Vitreisporites craigii*; 80. *Veryhachium formosus*; 81. *Leiofusa jurassica*; 82-83 *Luehndea spinosa*; 84. *Micrhystridium* sp.

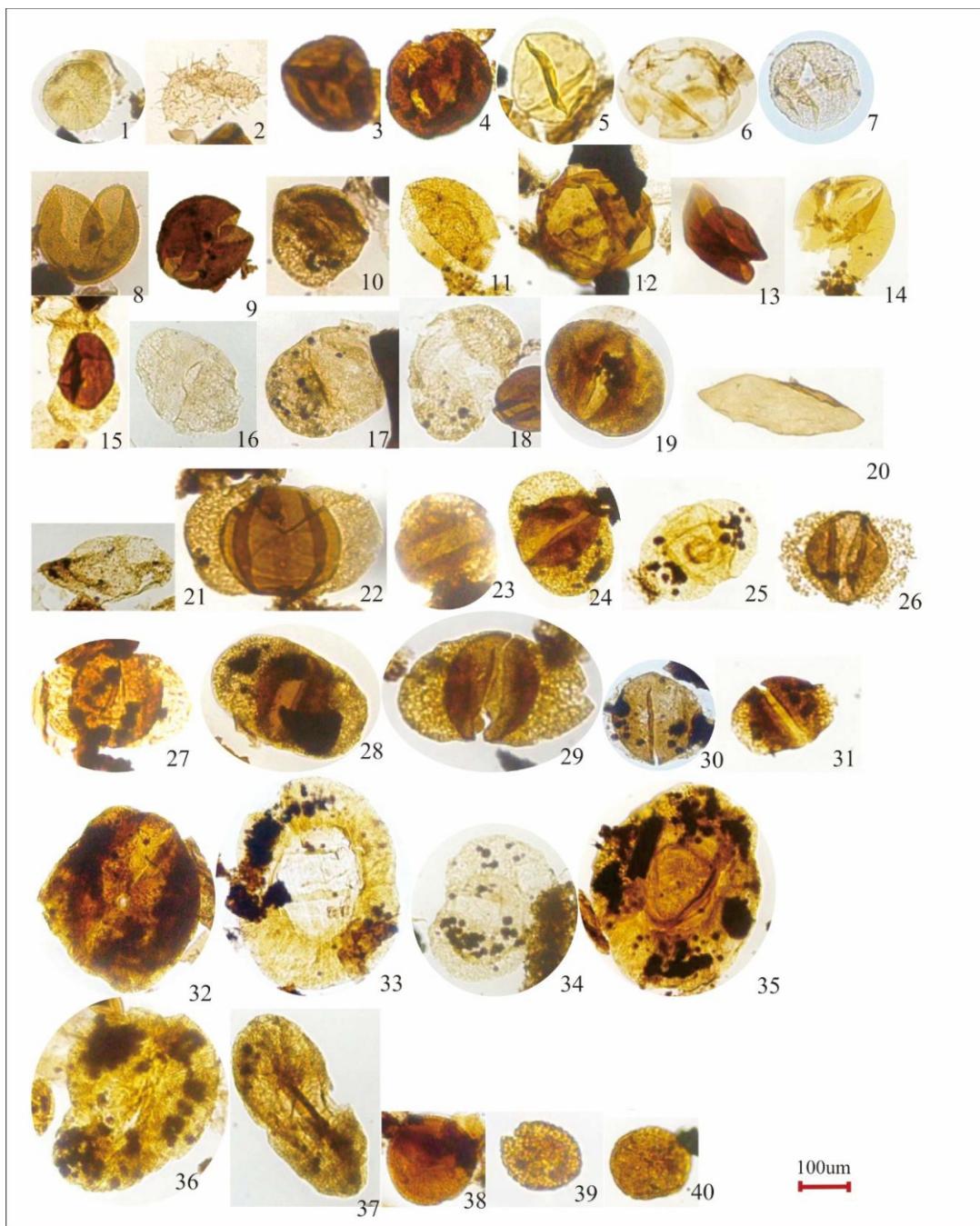


Fig. 7. Palynomorphs from the Lower Jurassic of the Indus Basin; 1. *Micrhystridium* sp.; 2. *Polygonium jurassicum*; 7. cf. *Chasmatosporites elegans*; 8-9. *Cythatides australis*; 10. cf. *Alisporites microsaccus*; 11. *Chasmatosporites major*; 12. *Araucariacites australis*; 13. *Spheripollenites* sp.; 14. *Cyclotriletes* sp.; 15. *Lunatisporites kraeuseli*; 16-19. *Vitreisporites pallidus*; 20. *Cycadopites dilucidus*; 21. *Dinocyst* indet.; 22-26. *Quadraeculina anellaeformis*; 27-28. *Alisporites pergrandis*; Fig. 29. *Verrucosporites* sp.; 30-32. *Dipterella oblatinoides*; 33-34. *Chasmatosporites hians*; 35-36. *Aratrisporites fischeri*; 37. *Monosulcites minimus*; 38-40 *Corollina torosus*.

4.2 Bulk Carbon and Oxygen Isotopic Study

The $\delta^{13}\text{C}$ in the Datta Formation shows the minimum and maximum values of -2‰ to -3‰ negative carbon isotope excursion, with some samples have not recorded any experimental data in laboratory analysis. Furthermore, the Datta Formation show the minimum value of -0.95‰ and maximum value of -3.11‰ (Fig. 8).

4.3 Organic contents

The carbon contents is find out by the total organic content ranges from the 0.2-4.25%.

5. Discussion

5.1 Palynostratigraphic Frame work of Lower Jurassic

Dinoflagellates, spores, and pollens are utilized for stratigraphy of the Jurassic rocks all-around the globe (Pocock, 1978; Kotova, 1983; Hoelstad, 1985; Koppelhus and Nielsen, 1994; Tiwari and Tripathi, 1995; Burger, 1996; Dong et al., 1998; Tiwari, 1999; Tripathi, 2001; Koppelhus and Gregers, 2003; Ziaja, 2006; Atawy, 2009; Racey and Goodall, 2009; Barrón et al., 2010; Butaina 2010; Schrank, 2010; Goryacheva, 2011; 2017 El Atfy et al., 2013; Peng et al., 2018). Based on the palynomorph assemblage encountered in the Datta Formation the Hettangian-Pliensbachian assemblage biozone (HPABZ) was identified.

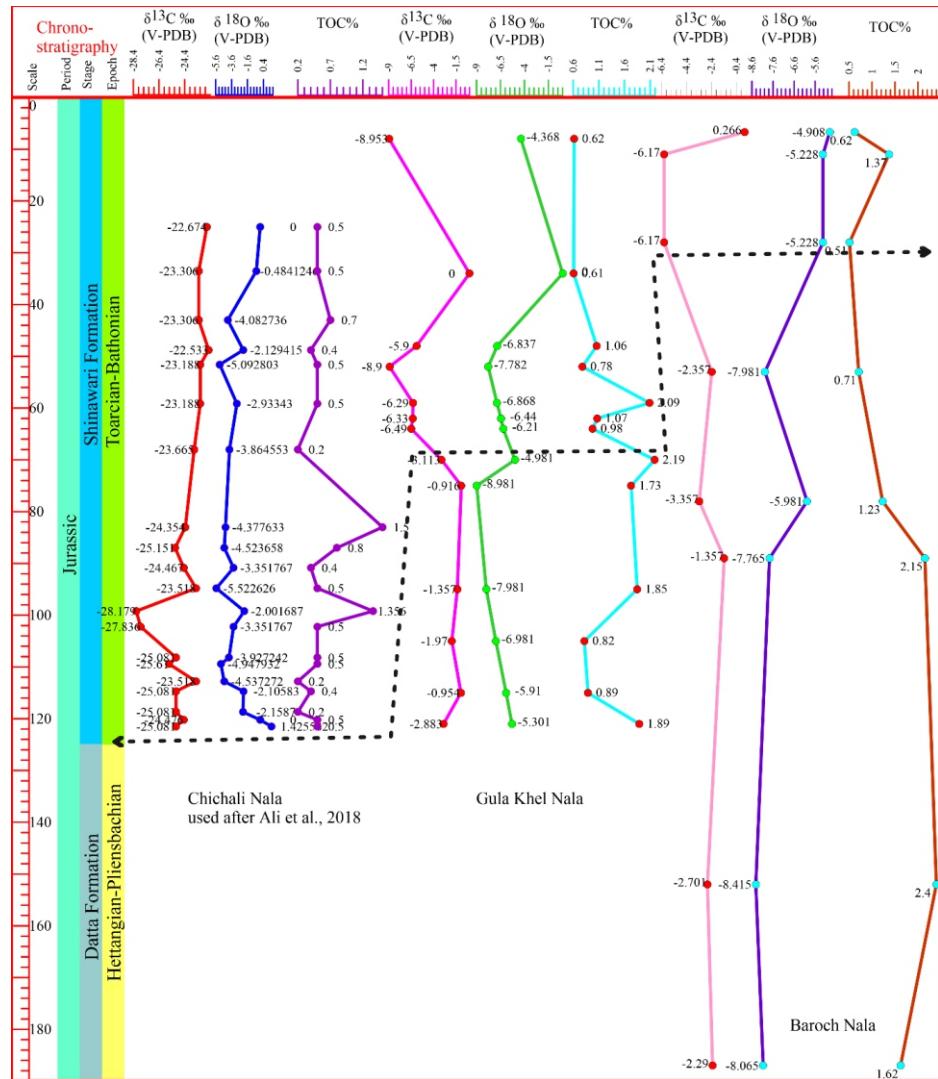


Fig. 8. Shows isotopic signals and TOC of Lower to middle Jurassic strata from Chichali Nala, Gula Khel Nala and Baroch Nala in the Indus Basin, Pakistan.

The presence of *Calamospora tener*, *Vitreisporites pallidus*, *Artrisporites minimus*, *Chasmatosporites* sp. such as *Chasmatosporites elegans*, *Chasmatosporites hians* and *Chasmatosporites major*, indicate Hettangian age (Schulz, 1967; Pocock, 1978; Pedersen and Lund, 1980; Lund and Pedersen, 1985; Dybkjær, 1991; Ziaja, 2006; Correia et al., 2018). The Sinemurian is represented by the *Corollina torosus*, *Vitreisporites pallidus*, *Corollina meyeriana*, and *Cerebropollenites macroverrucosus* (Bujak and Williams, 1977). The Hettangian to Pliensbachian age is considered for *Corollina torosus* biozone (Helby et al., 1987), whereas according to some other researchers a Pliensbachian age is suitable for the biozone represented by the *Chasmatosporites* (abundance of *C. hians* and *C. sp.*) and at least showing the presence of *Corollina torosus*. Moreover, the *Alisporites pergrandis*, *Camptotriletes cerebreformis*, *Chasmatosporites hians*, *Classopolis* sp., *Dipterella oblatinoides*, *Marattisporites scabratus*, *Osmundacidites* sp., *Stereisporites compactus*, *Stereisporites congregatus*, *Stereisporites bujargiensis*, *Stereisporites psilatus*, *Stereisporites infragranulatus*, *Vitreisporites pallidus* is similar to most of the Pliensbachian age Palynoassemblage 1 of Goryacheva (2017). So, based on the documented palynomorphs and the stratigraphic position in the Indus Basin, Pakistan, the lower Jurassic Datta Formation in the current study is marked as Hettangian to Pleinsbachian, while the Lower to Middle Jurassic Shinawari Formation is marked as Toarcian to Bajocain (Bender and Raza, 1995; Ali et al., 2018).

5.2 Anoxic Events

The TOAE was led by the PL-TO episode, with lesser $\delta^{13}\text{C}$ -2‰ to -3‰ negative carbon isotope excursion (Little and Benton, 1995; Pálfy and Smith, 2000; Hesselbo et al., 2007; Littler et al., 2010; Caruthers et al., 2013; Bodin et al., 2016; Percival et al., 2016). The Neotethys Tectonics of Hettangian-Pliensbachian strata in the current research is interpreted from the multi cyclic sequence representing the clastic and carbonates units (Fürsich et al., 1992). The large amplitude negative excursion in the bulk organic matter

and the similar trend in the $\delta^{13}\text{C}$ of the bulk carbonate suggests that the Toarcian strata in the Chichali Nala Section correspond to the TOAE (Ali et al., 2018). Furthermore, the Datta Formation show the minimum value of -0.95‰ and maximum value of -3.11‰ having well defined Pliensbachian-Toarcian boundary showing 2-3 ‰ per mil shift in the studied strata. Furthermore, the Nammal Gorge Section has no deposition of the PL-TO strata and onward while in Kahi Village the PL-TO strata are represented by carbonate rocks. Beside pronounced negative excursion, some positive values which may be due to one or a combination of various factors such as an increase in sedimentation rates and organic matter accumulation, marine anoxia or diagenetic basin evolution (Brasier et al., 1994). The TOC values of the dark black shale vary from 0.2-4.25% and more or less correspond to the $\delta^{13}\text{C}$ excursion in the studied area. A similar correlation between TOC and the carbon isotope excursion is reported by the previous researchers (Ali et al., 2018).

The current research has two folds importance i.e. the stratigraphic information can be utilized by the oil and gas sectors for formations/lithologies identifications and other related process. While, the academia can get benefit from the established stratigraphy and anoxia for further geological research.

6. Conclusions

The following conclusions are drawn;

1. The palynostratigraphy of the Lower Jurassic is established in the Datta Formation and Hettigian-Pliensbachians assemblage (HPABZ) biozone is identified. The Hettangian age indicators species are *Chasmatosporites elegans*, *Chasmatosporites hians* and *Chasmatosporites major*, while the Sinemurian is represented by the *Corollina torosus*, *Vitreisporites pallidus*, *Corollina meyeriana*, and *Cerebropollenites macroverrucosus* and Pliensbachian age is represented by the *Chasmatosporites*.
2. The Pliensbachian-Toarcian boundary is demarcated based on 2-3 ‰ shift in stable carbon and Oxygen isotopes in all the studied

sections except Nammal Gorge and Kahi village sections.

Authors' Contribution

Fahad Ali and Shiqi Zhang, proposed the main concept and involved in write up. Beenish Ali, assisted in establishing figure file. Rafique Ahamed, collected field data and formatting. Hamid Iqbal and Rafique Ahmad, did provision of relevant literature, and review and proof read of the manuscript.

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Appendices

Alisporites pergrandis, (Bolch.) Iljina
Aratrisporites fischeri, (Klaus) Playford and Dettmann, 1965
Aratrisporites minimus, (Schulz 1967)
Araucariacitesaustralis, (Cookson 1947)
Calamosporatener, (Leschik 1955) Mädler 1964
Camptotriletescerebreformis, (Naum. ex Jarosh)
Cerebropollenitesmacroverrucosus, (Thiergart) Schulz 1967
cf. *Alisporitesmicrosaccus*, (Couper 1958) Pocock 1962
cf. *Chasmatosporiteselegans*, (Nilsson 1958)
Chasmatosporiteshians, (Nilsson 1958)
Chasmatosporites major, (Nilsson 1958)
Pocock & Jansonius 1969
Chomotriletes minor, (Kedves) Pocock 1970
Classopollistorosus, (Reissinger 1950) Couper 1958
Clavatipolleniteshughesii, (Couper 1958)
Corollinameyeriana, (Klaus 1960)
Venkatachala & Goczán 1964
Cycadopitesdilucidus, (Bolch.) Iljina
Cythatidesaustralis, (Couper 1953)
Deltoidosporacrassexina, (Nilsson 1958) Lund 1977

Dipterellaoblatinoides, (Maljkina 1949)
Leiofusajurassica, (Schulz 1967)
Luehndeaspinosa, (Morgenroth 1970)
Lunatisporiteskraeuseli, (Leschik 1955)
Lycopodiumsporitescerniidites, (Ross 1949)
Delcourt & Sprumont 1955
Lycopodiumsporitesreticulumsporites, (Rouse 1959) Dettmann 1963
Marattisporitescabratus, (Couper 1958)
Monosulcitesminimus, (Cookson 1947) Couper 1953
Perinopolleniteselatoides, (Couper 1958)
Piceapollenitesvariabiliformis, (Bolch.) Petr., *Polygoniumjurassicum*, (Bucefalo Palliani et al., 1996)
Quadratraculinaanellaeformis, (Maljkina 1949)
Spheripollenitespsilatus, (Couper 1958)
Stereisporitesbujargiensis, (Bolch.) Schulz 1967
Stereisporitescompactus, (Bolch.) Iljina
Stereisporitescongregatus, (Bolch.) Schulz 1967
Stereisporitesinfragramulatus, (Schulz 1967)
Stereisporitespsilatus, (Ross) Pflug 1953
Todisporites minor, (Couper 1958)
Veryhachiumformosus, (Stockmans & Williere 1960)
Vitreisporitespallidus, (Reissinger 1950)